## Paper:

# Installation of New GNSS Network Around Kusatsu-Shirane Volcano, Japan: Its Perspective and the First Result 

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Crustal deformation is essential information for monitoring volcanic activity. In the summit area of the Kusatsu-Shirane Volcano (KSV), a dense Global Navigation Satellite System (GNSS) network has been operating near the recent volcanic center, Yugama crater. This network is sensitive to shallow depth activity, such as phreatic eruptions at the summit area, but is not applicable to deep magmatic activity, suggested to have been occurring for thousands of years by recent geological studies. Aiming to detect magmatic activity at a certain depth, we installed a new GNSS network near KSV. The observation sites were selected based on the crustal deformation pattern calculated for several intrusive events of the deep-seated magma. First, the GNSS sites for campaign observation were installed at eight locations in 2017. Then, four continuous sites commenced operation after a phreatic eruption at Mt. Motoshirane in January 2018. Here, we show the results of the first and second observation campaigns, operating in October 2017 and February 2018. Coordinate values are computed by precise point positioning with ambiguity resolution (PPPAR) analysis and are used to calculate the displacement and the baseline length change during this period. The uncertainties of the calculated coordinate values are sufficiently small (less than 4.5 mm ) except at some sites for which the data possibly include multipath errors due to trees and snow. Although any deformation associated with the 2018 eruption of Mt. Motoshirane is not detected, subsequent observations would contribute to monitoring long-term activity near KSV.

Keywords: Kusatsu-Shirane volcano, global navigation satellite system, crustal deformation, magma reservoir

## 1. Introduction

The Kusatsu-Shirane volcano (KSV) is one of the active volcanoes in Japan and is known to have a welldeveloped hydrothermal system which is closely related to phreatic eruptions and seismic activity (e.g., [1-3]). KSV consists of two major peaks: Mt. Shirane and Mt. Motoshirane (Fig. 1). Mt. Shirane has three major crater lakes (Yugama, Mizugama, and Karegama). Mt. Motoshirane is composed of four main pyroclastic cones: Ko-Motoshirane, Shin-Motoshirane, Kagami-ike, and Kagami-ike-kita ([4] and references therein). They are all volcanic vents, some of which have fumarolic activity. Between Mt. Shirane and Mt. Motoshirane, there exist several vents, including Ainomine pyroclastic cone and Yumi-ike maar. Yumi-ike maar is thought to have been formed simultaneously with pyroclastic cones in the Mt. Motoshirane area [5]. Recent eruption activity of KSV is represented by phreatic eruptions near Yugama crater at Mt. Shirane. Since the oldest historic eruption of 1882, nineteen phreatic eruptions have been recorded [6]. After the last magmatic eruption at Kagami-ike-kita cone nearly 1500 years ago, the volcanic activity in KSV shifted to the phreatic eruption activity near Yugama crater at Mt. Shirane (e.g., [4]). In other words, KSV is characterized by the more recent phreatic eruptions, which have exhibited continuous magmatic activity at Mt. Motoshirane since approximately 1500 years ago (e.g., $[4,7]$ ). Recent geological studies have revealed magmatic eruptions that caused large amounts of lava effusion during the past thousands of years in $\operatorname{KSV}[4,8,9]$. In this background, it is necessary to prepare an observation network to detect the crustal deformation in a wide range leading to a magmatic eruption in the future. For this purpose, we installed a new GNSS network around KSV.



Fig. 1. Newly installed and pre-existing GNSS network around Kusatsu-Shirane Volcano. Black dashed line box in the left side map indicates the region of the right side map. Base map: ALOS Global Digital Surface Model which is 30-m-mesh data, issued by JAXA.

## 2. Background: Pre-Existing Observation Sites

Since around 2009, the Kusatsu-Shirane Volcano Observatory (KSVO) of the Tokyo Institute of Technology has been studying crustal deformation related to the volcanic activity of KSV, using several observation instruments consisting of seismometers, tiltmeters, and a Global Navigation Satellite System (GNSS) network. Because the recent eruptive activity is characterized by phreatic eruptions in and around Yugama crater [5], the pre-existing observation network was constructed with the aim of detecting shallow crustal deformation near Yugama crater: KSR, KSW, KSE, and KSYG (added in 2011) (Fig. 1). Using this GNSS network, crustal deformation associated with volcanic unrest has been observed continuously. The collected data are sent to the data server in KSVO and automatically computed as the daily coordinate values by Spider software (Leica). Whereas this high-density network is sensitive to shallow activity, detection of deep crustal deformation caused by intrusions of magma can be limited. To detect deeper crustal activity, a wider-range observation network is necessary. Furthermore, this concentrated network is vulnerable to eruptions in the summit area. Actually, in the 2018 eruption of Mt. Motoshirane, the observation network faced a crisis of power failure as the commercial power system was destroyed by volcanic explosions. To avoid such a situation, it is meaningful to expand the observa-
tion network to locations away from the volcanic center. Besides KSVO, the Geospatial Information Authority of Japan (GSI), Japan Meteorological Agency (JMA), and National Research Institute for Earth Science and Disaster Resilience (NIED) also have continuous GNSS observation sites near KSV (Fig. 1). Although data from these other institutions, covering a wide range around KSV, can be used, the observation sites are sparsely distributed. In particular, there are few observation sites within $5-10 \mathrm{~km}$ from Yugama crater (Fig. 1), which are considered to be essential to detect the crustal deformation caused by the magma activity of KSV. Therefore, we have to increase the spatial density of the GNSS network.

## 3. Method

### 3.1. Selection of Observation Sites

To select the location of an observation site to be newly installed, we calculated the uplift height and horizontal displacement associated with a deep magmatic intrusion (modeled as a spherical inflation source) in advance of using the Mogi model [10]. Considering the topography of the volcanic edifice, we used the elevation-corrected Mogi model [11];

$$
\begin{equation*}
U_{z}=\frac{3 a^{3} \Delta P}{4 \mu} \frac{D+h}{\left\{(D+h)^{2}+r^{2}\right\}^{\frac{3}{2}}} \tag{1}
\end{equation*}
$$

$$
\begin{equation*}
U_{r}=\frac{3 a^{3} \Delta P}{4 \mu} \frac{r}{\left\{(D+h)^{2}+r^{2}\right\}^{\frac{3}{2}}} \tag{2}
\end{equation*}
$$

where $U_{z}$ and $U_{r}$ are the vertical and horizontal displacements, respectively, $a$ is the radius of the inflation source, $D$ is the depth of inflation source below sea level, $h$ is the height above sea level, $\Delta P$ is the pressure change, and $\mu$ is the shear modulus. We assumed 1000 m , $1.01325 \times 10^{8} \mathrm{~Pa}$, and 40 GPa for $a, \Delta P$, and $\mu$, respectively. These assumptions correspond to a volume increase of $7.85 \times 10^{6} \mathrm{~m}^{3}$. The horizontal and vertical displacements were calculated using MaGCAP-V ver. 1.7 (Magnetic and Geodetic Computer Analysis Processes for Volcano [12]) with varying depths for the inflation source in the range of $0-10 \mathrm{~km}$ [11]. Fig. 2 and Figs. 4-6 (in Appendix A) show examples of calculation results assuming the depth of the inflation source to be 2.5 km , $5.0 \mathrm{~km}, 7.0 \mathrm{~km}$, and 10.0 km below sea level just beneath Shin-Motoshirane crater or Yugama crater. These results indicate that the crustal deformation associated with deep magmatic intrusions cannot be detected at the pre-existing GNSS observation sites that are concentrated around Yugama crater; thus, it is necessary to disperse the observation sites within 15 km of the summit area. Considering the simulation results and accessibility, we determined the location of ten sites consisting of four continuous sites (KSH1, KSM1, KSZG, and SHG1) and six campaign sites (NHH, TMG, YNG, OGR, NZL, and NKD). Table ?? shows the locations of those sites and the detailed observation settings. As described later, KSM1 and SHG1 started continuous observation after conversion from campaign observation sites (MSN and SSG, respectively).

### 3.2. Installation

We installed GNSS observation sites from the summer of 2017 to the autumn of 2018. In 2017, eight sites (MSN, SSG, NHH, TMG, YNG, OGR, NZL, and NKD) were set up for campaign observation. In 2018, four continuous observation sites (including two conversions from the campaign sites) were installed. Six of the campaign observation sites were built using a stainless-steel plate, upon which a tribrach adapter was mounted (Fig. 7 in Appendix A). This stainless-steel plate was attached to a preexisting building using a rubber sheet so as to track five or more GNSS satellites. For MSN, located at the southern rim of Shin-Motoshirane crater, we fixed a stainless-steel bolt, 6 cm in length, onto a stiff rock constituting a lava dome (Fig. 7B in Appendix A). For TMG, we buried a pin on the asphalt as a marker for tripod installation (Fig. 7C in Appendix A). We carried out campaign observations during a calm and dry autumn season without typhoons to avoid a warm and wet climate. During the campaign observation, four sets of GNSS equipment were used to perform a three-day observation twice. Each set of GNSS equipment consists of an antenna (JAVAD GrAnt G3T) and a receiver (JAVAD Sigma G2T). The sampling rate and elevation mask were set at 30 s and $10^{\circ}$, respectively.

In conjunction with the installation of campaign observation sites, we set up the continuous sites around KSV one after another: Shizukayama (KSZG) in March 2018 and Manza (KSH1) in May 2018 (Fig. 1). The former is located at approximately 2.5 km WNW from the Yugama crater, and the latter is approximately 5.5 km SE from the crater. In addition, we converted a campaign site (SSG) into the continuous one (SHG1) in May 2018. At these three sites, we installed the GR10 GNSS receiver and AR10 antenna (Leica Geosystems). AR10 is a multipurpose antenna with an integrated radome, which could prevent ash from adhering to the antenna. Finally, we started operating the campaign site (MSN) as a continuous site (KSM1) using the same GNSS equipment used for campaign observation (JAVAD Sigma G2T and GrAnt G3T) in October 2018. The data measured at these continuous observation sites are telemetered using the mobile phone network (Docomo 3G FOMA). Specifically, the $\mathrm{Wi}-\mathrm{Fi}$ router is turned on by a mechanical timer for only one hour (9:00-10:00 JST) every day to establish an internet connection for the GNSS receiver. During that time, the raw carrier phase and pseudo-range data stored in the receiver are downloaded via FTP from the data server in KSVO. This extension of the GNSS observation network is expected to lead to the detection of a wider range of crustal deformation, caused by inflation of the deep magma reservoir of KSV.

### 3.3. Data Processing

For data measured by the campaign observation, we performed precise point positioning with ambiguity resolution (PPP-AR) [13] analysis to determine the position of each site; then, we calculated their displacements and the baselines between them. PPP-AR is a positioning technique that is able to compute an accurate solution with a single receiver without any reference stations [13]. Because the raw data were recorded as a JPS file (native format file of JAVAD GNSS receiver), they were converted to a RINEX file using JPS2RIN software (ver. 2.0.112, JAVAD GNSS). Then, the converted RINEX data were used to calculate daily positions in reference to the IGS08 by PPP-AR analysis using GIPSY OASIS II ver. 6.1.2, developed at the NASA Jet Propulsion Laboratory (JPL) [13]. Taking into account the ocean tide loading values, FES2004 was adopted as ocean tide model for each observation site [14]. During the data processing, we also considered the GNSS antenna phase characteristics consisting of antenna phase center offset and phase center variation factor of the antenna. The wet atmospheric delay effect was corrected using the global mapping function (GMF) [15], based on numerical weather model data obtained by JPL. The clock data files were downloaded from the FTP server of the Crustal Dynamics Data Information System (CDDIS), provided by NASA. Calculation of relative displacements and mapping of their spatial distribution were also performed using MaGCAP-V. We adopted the GEONET 020982 site, one of the nationwide GNSS network sites operated by GSI, as the reference site


Fig．2．Distribution of horizontal displacement（red arrow）based on the Mogi model when an inflation source is assumed to be located beneath the Shin－Motoshirane crater at $0,2.5,5.0,7.5$ ，and 10 km below sea level．

Table 1．Description of GNSS stations around Kusatsu－Shirane volcano including pre－existing sites of KSVO．

| Site | Place name | in Japanese | Latitude［degree］ | Longitude［degree］ | Elevation $[\mathrm{m}]$ | Continuous／Repeated | Power source |
| :--- | :--- | :--- | :--- | :--- | :--- | :--- | :--- |
| KSZG | Shizukayama | 静可山 | 36.6010 | 138.5721 | 1270 | Continuous | AC power |
| KSH1 | Manza | 万座 | 36.6344 | 138.5100 | 1681 | Continuous | AC power |
| NHH | Naganohara | 長野原 | 36.5514 | 138.6317 | 624 | Repeated | Battery |
| TMG | Tsumagoi | 嬬恋 | 36.5481 | 138.4654 | 1394 | Repeated | Battery |
| YNG | Yonago | 米子 | 36.6123 | 138.3577 | 690 | Repeated | Battery |
| NKD | Kasadake | ニュー笠岳 | 36.6670 | 138.4603 | 1492 | Repeated | Battery |
| SSG／SHG1 | Shiga Kogen | 志賀高原 | 36.7108 | 138.4951 | 1595 | Repeated $\rightarrow$ Continuous | Solar＋Battery |
| NZL | Nozoriko | 野反湖 | 36.6914 | 138.6520 | 1591 | Repeated | Battery |
| OGR | Ogura | 小倉 | 36.6656 | 138.6115 | 1150 | Repeated | Battery |
| MSN／KSM1 | Motoshirane | 本白根 | 36.6212 | 138.5347 | 2135 | Repeated $\rightarrow$ Continuous | Solar＋Battery |
| KSE | Yugama higashi | 湯釜東 | 36.6438 | 138.5446 | 1941 | Continuous | AC power |
| KSYG | Yugama | 湯釜 | 36.6413 | 138.5347 | 2057 | Continuous | AC power |
| KSR | Rest house | レストハウス | 36.6378 | 138.5344 | 2012 | Continuous | AC power |
| KSW | Yugama nishi | 湯釜西 | 36.6471 | 138.5267 | 2096 | Continuous | AC power |

${ }^{1}$ Stopped during October 2018 and November 2018.
to calculate the relative displacement so as to remove the common crustal deformation caused mainly by the post－ seismic deformation of the 2011 Tohoku－Oki earthquake （e．g．，［16］）．

To calculate the baseline lengths at each site，we used calculation services provided by GSI（https：／／vldb．gsi．go． jp／sokuchi／surveycalc／main．html）for two steps of con－ version of the coordinate values 1）from earth－centered，
earth－fixed（ECEF）to latitude－longitude－geoid height， and 2）from latitude－longitude－geoid height to north－east－ up（NEU）．For the NEU reference frame，we applied the Japan Plane Rectangular CS（VIII for SHG，NKD，YNG， and IX for NHH，OGR，NZL，TMG，KSM，KSYG，KSR， and KSW）．Baseline lengths were calculated using the Pythagorean theorem in three－dimensions．


Fig. 3. Displacements during 2017 October to 2018 February around Kusatsu-Shirane Volcano. Dashed rectangle in (A) indicates the region of (B).

## 4. Results: The First and Second Campaign Observations

The first campaign observation was performed from middle to late October 2017. After the occurrence of a phreatic eruption at Mt. Motoshirane on January 23, 2018 [3], we performed an urgent temporary observation during early February. On account of snow coverage, some sites were inaccessible; eventually, we were able to access only five sites: NHH, YNG, OGR, NKD, and SSG in the latter observation. As described above, we carried out the measurements for more than three days at each site; thus, we analyzed the daily data. Therefore, the mean coordinate values were adopted as the positioning results for the observation period, in which apparent outliers were excluded. Table 4 (in Appendix A) shows the coordinate values of each GNSS observation site, including pre-existing sites of KSVO.

The standard deviations, representing the accuracy of the first campaign observation, are sufficiently small (less than 5.3 mm and 21.7 mm for horizontal and vertical directions, respectively). Given the surrounding circumstances, a multipath error caused by nearby trees is possibly contained in the data at OGR. For the second campaign observation, it is considered that the OGR and SSG data contain multipath errors caused by the surrounding snow.

Figure 3 and Table 2 show the displacement of each site and the change in baseline length (distance between each site with Naganosakae) between the first campaign observation and the second one. During October 2017 to February 2018, the horizontal displacement shows an eastward trend at all sites, except YNG and OGR. The vertical displacements are positive (i.e., upward) for

Table 2. The amount of displacement relative to the electronic reference point of GSI (Naganosakae, 020982) during October 2017 to February 2018 detected by our GNSS network.

| Site | Eastward <br> displace- <br> ment [mm] | Northward <br> displace- <br> ment [mm] | Upward <br> displace- <br> ment [mm] $]$ | Baseline <br> change <br> between <br> Naganosakae <br> $[\mathrm{mm}]$ |
| :--- | :--- | :--- | :--- | :--- |
| NHH | 2.3 | -0.9 | 21.5 | 0.8 |
| OGR | -14.3 | 0 | -3.1 | 0.7 |
| SSG | 0.7 | 5.3 | -56.6 | -6.6 |
| NKD | 1.3 | 3.1 | 24 | -2.8 |
| YNG | -0.1 | 5 | -30.5 | -3.5 |
| KSYG | 4.1 | 7 | 5.4 | -7.7 |
| KSW | 1.7 | 11.3 | 11.5 | -10.7 |
| KSR | 7.3 | 6.7 | 5.4 | -8.4 |

NHH, NKD, KSYG, KSW, and KSR, and negative (i.e., downward) for OGR, SSG, and YNG. Table 3 shows changes in the length of each baseline (between each pair, XYZ in reference to the GRS80) during October 2017 to February 2018. The baseline change in this period is less than 18.2 mm . Because our data is not continuous (i.e., campaign observation for a few days) and the scale of the 2018 eruption at Mt. Motoshirane was small, it is nearly impossible to discuss the contribution of the eruptive activity to our data.

Table 3. Change in baseline length between each site during October 2017 to February 2018 detected by newly built GNSS network.

|  | Baseline length change [mm] |  |  |  |  |  |  |  |
| :--- | :--- | :--- | :--- | :--- | :--- | :--- | :--- | :--- |
|  | NHH | NKD | OGR | SSG | YNG | KSYG | KSR | KSW |
| NHH | - | -1.3 | 4.7 | 1.0 | -6.0 | 5.8 | 1.9 | 11.4 |
| NKD | -1.3 | - | 10.6 | -2.1 | 8.1 | -2.5 | -10.8 | -4.8 |
| OGR | 4.7 | 10.6 | - | 13.5 | 6.7 | -10.4 | -18.2 | -12.8 |
| SSG | 1.0 | -2.1 | 13.5 | - | 4.2 | 0.6 | -7.7 | -1.9 |
| YNG | -6.0 | 8.1 | 6.7 | 4.2 | - | -5.9 | -14.0 | -7.6 |
| KSYG | 5.8 | -2.5 | -10.4 | 0.6 | -5.9 | - | -2.7 | 4.0 |
| KSR | 1.9 | -10.8 | -18.2 | -7.7 | -14.0 | -2.7 | - | 8.9 |
| KSW | 11.4 | -4.8 | -12.8 | -1.9 | -7.6 | 4.0 | 8.9 | - |

## 5. Summary and Perspective

We installed GNSS observation sites near KusatsuShirane Volcano in an effort to detect the crustal deformation associated with deep magma intrusion. The first observation result shows that our extended GNSS network works well despite some exceptions due to surrounding circumstances. In the present study, crustal deformation associated with the phreatic eruption at Mt. Motoshirane was not detected. The primary factor is its smallness in eruption scale; however, we also have problems that need improving in terms of analysis. To extract the crustal deformation caused by the magmatic activity of KSV, several influences, such as the 2011 earthquake off the Pacific coast of Tohoku have to be corrected from the measured positioning data. We believe that the continuation of observations would contribute to the monitoring of longterm volcanic activity near Kusatsu-Shirane Volcano.

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## Appendix A. Supplementary Figures and Table

Figures 4-6 present the calculation results based on Mogi model and Fig. 7 shows observation sites. Table 4 shows the coordinate values of each GNSS observation site.


Fig. 4. Distribution of vertical displacement (red arrow) based on the Mogi model when an inflation source is assumed to be located beneath the Shin-Motoshirane crater at $0,2.5,5.0,7.5$, and 10 km below sea level.


Fig. 5. Distribution of horizontal displacement (red arrow) based on the Mogi model when an inflation source is assumed to be located beneath the Yugama crater at $0,2.5,5.0,7.5$, and 10 km below sea level.


Fig. 6. Distribution of vertical displacement (red arrow) based on the Mogi model when an inflation source is assumed to be located beneath the Yugama crater at $0,2.5,5.0,7.5$, and 10 km below sea level.


Fig. 7. Photos of observation sites. A: NHH, B: MSN, and C: TMG.
Table 4. Coordinate values by newly built GNSS campaign network.

| Cooxdinate values. NEU GRS.80 (vilux) (northeastup) |  |  |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: |
| $\times[m]$ | $Y$ [m] | 2 [m] | sad [m] | $\operatorname{sov}[m]$ | scz[m] |
| 61899.5968 | -10757.8873 | 677.9806 | 0.0007 | 0.0016 | 0.0076 |
| 61899.5951 | -107578.8570 | 677.9727 | 0.0008 | 0.0031 | 119 |
| 74541.6141 | ${ }^{109223.2676}$ | ${ }^{1194.4213}$ | 0.0007 | ${ }^{0.0018}$ | 0.015 |
| 74541.6688 | -102232.2705 | 1194.3780 | 0.0023 | 0.0005 | 0.0160 |
| ${ }^{77361.9333}$ | -10557.1211 | 1635.5612 | 0.0033 | 0.0027 | 0.0080 |
| 61887,0450 | ${ }^{122471.2556}$ | 1438.959 | 0.0013 | 0.0018 | 0.0054 |
| 78865.2617 | -435.5426 | 1647.1468 | 0.0017 | 0.0013 | 0.0018 |
| ${ }_{78665.2693}$ | -435.5321 | 1647.0727 | 0.0050 | 0.0029 | 217 |
| 74001.7 .7158 | .3550.7829 | 1540.3230 | 0.0045 | 0.0019 | 0.0065 |
| 74011.7261 | -3550.7753 | 1540.324 | 0.0025 | 0.0016 | 0.0114 |
| 67950.0784 | -12730.8188 | 730.5540 | 0.0013 | 0.0017 | ${ }^{0.0057}$ |
| 67950.032 | -12730.8148 | 730.9273 | 0.0020 | 0.0035 | 0.0172 |
| 69704.457 | -116157.489 | 2188.1628 | 0.0017 | 0.0028 | 0.0079 |
| 71945.1067 | -116122.5442 | 2099.7542 | 0.0018 | 0.0017 | 0.0172 |
| 71945.1137 | -116122.5440 | 2099.7374 | 0.0015 | 0.0009 | 0.0050 |
| 71556.4718 | -116157.6436 | 2058.7072 | 0.0021 | 0.0037 | 0.0147 |
| 71556.4803 | -116157.6252 | 2055.9948 | 0.0011 | 0.0013 | 0.0044 |
| ${ }^{72597.5213}$ | -1168324765 | 21412168 | 0.0011 | 0.0053 | 0.0077 |
| ${ }^{72597.5358}$ | -116832.4642 | 2141.2136 | 0.0008 | 0.0011 | 0.00 |








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Researcher, Institute of Space and Astronautical Science, Japan Aerospace Exploration Agency

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2015- Institute of Space and Astronautical Science, Japan Aerospace Exploration Agency
2015- Geological Survey of Japan, National Institute of Advanced Industrial Science and Technology
2016- Volcanic Fluid Research Center, School of Science, Tokyo Institute of Technology
2018- Institute of Space and Astronautical Science, Japan Aerospace Exploration Agency

## Selected Publications:

- R. Noguchi and K. Kurita, "Unique characteristics of cones in Central Elysium Planitia, Mars," Planetary and Space Science, Vol.111, pp. 44-54, 2015.
- R. Noguchi, A. Höskuldsson, and K. Kurita, "Detailed topographical, distributional, and material analyses of rootless cones in Myvatn,
Iceland,'J. of Volcanology and Geothermal Research, Vol.318, pp. 89-102, 2016.
- D. Shoji, R. Noguchi, S. Otsuki, and H. Hino, "Classification of volcanic ash particles using a convolutional neural network and probability," Nature, Scientific Reports, Vol.8, Article No.8111, 2018.
Academic Societies \& Scientific Organizations:
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## Selected Publications:

- T. Nishizawa, H. Nakamura, T. Churikova, B. Gordeychik, O. Ishizuka,
S. Haraguchi, T. Miyazaki, B. S. Vaglarov, Q. Chang, M. Hamada, J.-I. Kimura, K. Ueki, C. Toyama, A. Nakao, and H. Iwamori, "Genesis of ultra-high-Ni olivine in high- Mg andesite lava triggered by seamount subduction," Nature, Scientific Reports, Vol.7, Article No. 11515 doi:10.1038/s41598-017-10276-3, 2017.
Academic Societies \& Scientific Organizations:
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## Brief Career:

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2010- Associate Professor, Volcanic Fluid Research Center, Tokyo Institute of Technology
2016- Associate Professor, School of Science, Tokyo Institute of Technology

## Selected Publications:

- W. Kanda et al., "A Heating Process of Kuchi-erabu-jima Volcano, Japan, as Inferred from Geomagnetic Field Variations and Electrical Structure," J. Volcanol. Geotherm. Res., Vol.189, No.1-2, pp. 158-171, 2010.
- W. Kanda et al., "Hydrothermal system of the active crater of Aso volcano (Japan) inferred from a three-dimensional resistivity structure model," Earth, Planets and Space, Vol.71, No.1, Article: 37, 2019.


## Academic Societies \& Scientific Organizations:

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- Society of Geomagnetism and Earth, Planetary and Space Sciences (SGEPSS)
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2006- Aso Volcanological Laboratory, Graduate School of Science, Kyoto University
2009- Volcanic Fluid Research Center, School of Science, Tokyo Institute of Technology

## Selected Publications:

- A. Terada and Y. Ida, "Kinematic features of isolated volcanic clouds revealed by video records," Geophys. Res. Lett., Vol.34, No.1, L01305, doi:10.1029/2006GL026827, 2007.
- A. Terada, T. Hashimoto, and T. Kagiyama, "A water flow model of the active crater lake at Aso volcano, Japan: fluctuations of magmatic gas and groundwater fluxes from the underlying hydrothermal system," Bulletin of Volcanology, Vol.74, No.3, pp. 641-655, 2012.
- A. Terada and T. Hashimoto, "Variety and sustainability of volcanic lakes: Response to subaqueous thermal activity predicted by numerical model," J. Geophys. Res., Solid Earth, Vol.122, No.8, pp. 6108-6130, 2017.

Academic Societies \& Scientific Organizations:

- Volcanological Society of Japan (VSJ)
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